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Hydrological Studies of Artal Sub-watershed of Belagavi District, Karnataka, India

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Authors' contributions

This work was carried out in collaboration among all authors. All authors read and approved the final manuscript.

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ABSTRACT

Watershed is also a hydrological response unit and a holistic ecosystem in terms of the materials, energy and information present. The Artal sub-watershed (4d5e3b) is located in the Athani taluk of Belagavi district. It lies between 16°43'50" - 16°45'40" North latitudes and 75°17'30" - 75°23'0" East longitudes and covers an area of about 4632.04 ha. The average annual rainfall (2016 - 2021) of Artal sub-watershed was 536.0 mm. The maximum precipitation of 357.3 mm is received from June to September, 108.9 mm from October to early December and the remaining 99.8 mm is received

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during the rest of the year. The number of rainy days (>2.5 mm) varied from 14-24 days per year. On average, number of rainy-day events likely to produce runoff (20 to 30 mm) is about 3 to 7 rainy days per year with moderate variation across the years. The Actual Evapotranspiration (AET) over the years 2016 - 2021 in the Artal sub-watershed varied from 297.0 to 558.1 mm. During 2016 - 2021, the average annual AET (421.0 mm) was less than the average rainfall (536.0 mm). The average AET/P ratio between 2016-2021 was about 0.74 which is less than the sustainable limit of about 0.80. During 2018 and 2020, the AET/P ratio were 1.14 and 1.06, respectively, which indicates receipt of less rainfall in these years and also a possibility of groundwater being augmented to maintain crop water requirements. A maximum area of about 875 ha (18.9 %) requires graded bunding and 3657 ha (78.96 %) area requires contour bunding.

Keywords: Watershed; evapotranspiration; groundwater; kriging; soil conservation measures.

1. INTRODUCTION

"The backbone of the Indian economy is agriculture and allied fields which mostly depend upon an abundance of natural resources like rainfall, water resources, soil and forest vegetation. As these resources are limited and depleted yearly, there is an utmost need to stabilize and conserve these resources. Watershed is the geographical area which is drained by the network of streams to the common outlet. A watershed is a complex and dynamic biophysical system which is identified as a planning and management unit. A watershed is also a hydrological response unit and a holistic ecosystem in terms of the materials, energy and information present. Land and water are the two main resources of the watershed. The watershed not only is a useful unit for physical analyses, it can also be an appropriate socio-economicpolitical component for the execution of management strategies. In essence, a watershed is a basic organizing unit to manage resources" [1].

"Unplanned and uncontrolled use of resources results in the deterioration of watershed which is further aggravated by human interventions. The population residing in the watershed is one of the most important assets of that watershed and the condition of a particular watershed depends heavily on its population. So, it is imperative to involve the people in conservational strategies actively involving proper management and execution of the watershed resources. The planning and management of the watershed are done to accomplish the tasks related to the overall development of the watershed, which may be with respect of water quality and quantity improvement, management of ecosystem, enrichment of the socio-economic status of the enhancing watershed inhabitants, the employment opportunity for the people and selection of most appropriate cropping pattern" [2].

"Watershed management is the balanced exploitation of terrestrial and aquatic resources for the acquisition of optimal production with petite vulnerability to natural assets. It adopts the rehearsal of soil and water conservational strategies in the watershed, for example, appropriate exploitation of the land, defensive measures of land against anthropogenic pressures, enhancement and management of soil fertility, water conservation for irrigational practices, proper supervision of local water supplies for drainage, protection against flash floods and reduction in runoff and soil erosion, and also escalating the production from all the existing land use patterns. Watershed-level hydrological studies are essential to soil and water resources evaluation, improvement and management. At the field scale, hydrologic studies are engaged in devising and designing of soil conservation practices" [3], management of irrigation water, water quality evaluation and water supply availability etc, [3]. Hydrological studies are important tools for comprehending the hydrological behaviour of the watersheds. Hence the hydrological studies have been taken up at the Artal sub-watershed (4d5e3b), which is located in the Athani taluk of Belagavi district, Karnataka.

2. METHODOLOGY

2.1 Location and Extent

The Artal sub-watershed is located in the south-eastern part of Karnataka in Athani taluk and Belagavi district (Fig. 1). It lies between $16^{0}43'50" - 16^{0}45'40"$ North latitudes and $75^{0}17'30" - 75^{0}23'0"$ East longitudes and covers an area of about 4632.04 ha. It is about 30 km from Athani town and is bounded by Halahalli on the east, Savalagi & Thungala on the south, Badagi & Aigali on the west and Thelasanga on the northern side of the sub-watershed.



Fig. 1. Location map of Artal Sub-watershed

2.2 Geology and Physiography of the Belagavi District

"Complex geological formations can be observed in the Belagavi district. The Schist and Banded ferruginous quartzite, the peninsular gneiss by granite and gneissic granites, the Kaladgi formations, sandstone, quartzite, shale and limestone and dolomite, basalt (Deccan Trap) and the laterite formations are observed in the district. The Deccan basalt occupies large extent in the Northern part, thinning out towards the South and it can be seen in Belagavi, Khanapur, Hukkeri, Chikkodi, Athani and Raibag taluks.The Khanapur taluk is enriched with a variety of fire clay, clastic clay, china clay etc" [3].

The district is divided into three physiographical divisions namely as Malenaadu Tract (Western Ghat Region), Gadinaadu Tract (Border area Region) and Bayalunaadu Tract (Plain Land Region). The "Malenaadu" tract is the Western Ghat area, with lush green forests, sharply undulating topography, and heavy rainfall. Many 1st order streams traverse this area. There are many natural springs in this tract. The "Gadinaadu" (intermediary) tract shows medium range flat to gently rising hills, with shrubby greenery, receiving an average rainfall. The streams are of 3rd & 4th order. The "Bayalunaadu" tract shows vast, flat terrain, with flat-topped barren hills. The rainfall received is less than 650 mm. The streams are very gentle.

2.3 Drainage, Rainfall and Ground Water Measurement

"The district falls under the influence of the Krishna River basin, and a small part of Khanapur taluka under the Kali River basin. The major tributary rivers of Krishna River are Ghataprabha and Malaprabha, which are supported by sub-tributaries like Markandeya, Hiranyakeshi, Doodhaganga, Vedganga, Agrani etc. The Malaprabha and Mahadai River originating in the Western Ghat of Khanapur taluka having several perennial springs, all along their courses. The Pandhari and Mahadai rivers in the south join the Kaali River basin" [2].

"Groundwater recharge is an important and required necessary activity in managing and developing water resources of a watershed. Rainfall is the major source of groundwater recharge. The infiltration of water is mainly governed by lithology, land use practice and elevation of the terrain. Spatial maps were prepared using ArcGIS 10.8. Month-wise data of the water table was collected in the artal sub watershed. The seasonal water-level fluctuations in the Artal watershed have been analysed. Rainfall data are one of the important datasets in the spatial domain, controlling the water resources budget of the region. Rainfall data for the last 06 years were collected from Karnataka State Natural Disaster Monitoring Centre, Govt of Karnataka" [3].

2.4 Point Interpolation: Kriging

"Kriging is a geostatistical method for estimating values in unknown areas by considering both the distance and variation between known data points. It involves creating an estimated surface from scattered points with z-values by fitting a mathematical function to nearby points. The process includes statistical analysis, variogram modelling, surface creation, and variance exploration. Predicted values are calculated using a weighted average technique based on the relationship between samples. The search radius can be fixed or variable and generated cell values may exceed the sample range" [4].

$$Z(S_0) \stackrel{N}{=} \sum_{i=1}^N \lambda Z(S_i)$$

Where,

Z(Si) = the measured value at the i th location

 λ_i = an unknown weight for the measured value at the i th location

 S_0 = the prediction location

N = the number of measured values

Kriging method is an interpolation "The method based on principles of zero bias and minimum mean square error. It determines values for process an entire а over domain, finite-volume block or specific point using a linear combination of data values. The summation may be over an entire area or restricted region centred at the estimation point" [5].

3. RESULTS AND DISCUSSION

3.1 Rainfall Analysis and Distribution

To develop the rainfall indices of the Artal subwatershed, data from the Artal rain gauge station in Athani taluk of Belagavi district was taken into account. During the year 2018, annual rainfall was deficient by 30 %. The district falls under the hot semiarid tract of the state and is categorized as drought -prone with an average annual rainfall of 565.9 mm received in the last 7 years (Table 1 and Fig. 2). A maximum of 357.3 mm of precipitation is received during south-west monsoon period from June to September, the north-east monsoon contributes about 108.9 mm and prevails from October to early December and the remaining 99.8 mm are received during the rest of the year. The winter season is from December to February. During April and May, the temperatures reach up to 38.8°C and in December and January, the temperatures will go down to 15.1°C. Rainfall distribution is shown in The average monthly Potential 3. Fig. Evapotranspiration (PET) is 139.1 mm and varies from a low of 102.4 mm in December to 191.7 mm in March.

Fable 1. Mean month	ly rainfall, PET	, 0.5 PET of Athan	i Taluk, Belagavi District
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Month	Rainfall (mm)	Temp- Max (°C)	Temp- Min (°C)	Max- RH (%)	Min- RH (%)	PET (mm)	0.5 PET (mm)
Jan	0.0	30.8	15.1	84.8	36.8	102.9	51.5
Feb.	0.0	33.1	16.6	75.1	28.0	121.3	60.7
Mar	8.9	36.7	19.9	72.6	22.1	191.7	95.9
Apr	29.8	38.8	22.5	80.7	22.7	171.5	85.8
May	61.2	37.9	23.6	88.1	30.7	177.1	88.6
Jun	76.9	33.1	22.8	94.5	54.5	146.9	73.5
Jul	93.8	30.0	22.2	95.4	65.7	151.8	75.9
Aug	56.9	29.3	21.7	96.9	66.8	146.7	73.4
Sep	129.7	30.3	21.3	97.2	60.9	114.1	57.1
Oct	93.9	31.2	20.6	93.8	51.0	139.1	69.6
Nov	11.2	31.1	18.5	87.5	44.2	103.9	52.0
Dec	3.8	30.2	16.3	87.9	43.1	102.4	51.2
Total	565.9					1669.4	834.7



Fig. 2. Average annual rainfall at Athani Taluk, Belagavi District





The PET is higher precipitation than months all except September. in the Generally, the length of the growing period in the Artal sub-watershed ranged from 150 to 160 days. The length of the growing period begins at 21st week (which is May 2nd week)) and ends at 43rd week (which is October 3rd week). Based on the observation, farmers can schedule sowing and other agronomic practices for short duration and long duration crops.

3.2 Rainfall Distribution

The number of rainy days (>2.5 mm) varied from 14-24 days per year. On average the number of rainy-day events likely to produce runoff (20 to 30 mm) is about 3 to 7 rainy days per year with moderate variation across the years (Fig. 4). The extremely high rainfall peak event days (>30mm) 3 day per year observed in the year 2017. Which helps in designing conservation and harvesting structure in a watershed.



Fig. 4. Distribution of Rainfall in different rainfall events



Fig. 5. Actual evapotranspiration (AET) and rainfall over different years

3.3 Evapotranspiration

The Actual Evapotranspiration (AET) over the years 2016-21 in the Artal sub-watershed varied from 297 to 558 mm (Fig. 5) During 2016-21 the average annual AET (421 mm) was less than the average rainfall (536 mm). AET was higher than rainfall during the months of January, February, March, April, August, November and December, which forced groundwater withdrawal during those months to meet crop water needs.

3.4 Annual Actual Evapotranspiration

The spatial maps of reference and actual evapotranspiration describe that the difference between ET_{ref} and ET_a is less over upper, middle and lower ridges of the watershed (Fig. 6). "The ridge-wise analysis in annual actual evapotranspiration shows that the extremely low (348-366 mm) values observed over upper ridge of the watershed and high (499-510 mm) amount of ET_a observed over lower ridge of the

watershed. Spatial variability in of actual evapotranspiration was due to soil moisture, vegetation cover and cropping systems adopted in the watershed" [6]. "To understand the reason for the variability in ET_a, the seasonal trends are calculated for the major evapotranspiration components including transpiration, bare soil evaporation, interception loss, and open water evaporation during the study period" [7]. The transpiration and interception loss from the vegetation has shown a tremendous change in the evapotranspiration [8].

3.5 Evapotranspiration Index

Budyko curve is the relationship between the ratio of the actual evaporation amount to the annual precipitation (AET/P) and the ratio of the potential evaporation amount to the annual precipitation (PET/P), called the dryness index or aridity index) based on data obtained from watersheds [9]. The watershed water balance is in normal condition. For sustainability, the limit of AET/P should be below the Budyko curve for sustainable watersheds from hydrological considerations. This suggests that the cropping

choices and irrigation choices have to be altered to reduce the total ET (Fig. 7). The average AET/P ratio between 2016-2021 was about 0.74 which is less than the sustainable limit of about 0.80. During 2018 and 2020 AET/P ratios were 1.14 and 1.06 respectively, (Fig. 8) which indicates receipt of less rainfall in the years and also the possibility of groundwater being augmented to maintain crop water requirement. To build on this and link the Budyko's hypothesis to the complementary relationship between actual evaporation and potential evaporation [10]. The Budyko curve analytically by modelling total evaporation using simple models of interception and transpiration in combination with measurable parameters related to rainfall dynamics and storage availability obtained from remotely sensed data sources [11]. "The Budyko curve has also been interpreted at a higher physical level as a possible outcome of thermodynamic optimality through the invocation of the maximum entropy production principle" [12]. "Budyko curve parameter could facilitate and easily utilized by policymakers for watershed quality assessment and hydrological system identification" [13].



Fig. 6. Map of the Annual Actual Evapotranspiration of Artal sub-watershed.

3.6 Soil Texture and Infiltration Rate

Texture is an expression to indicate the coarseness or fineness of the soil as determined by the relative proportion of primary particles of sand, silt and clay. It has a direct bearing on the structure, porosity, water infiltration, adhesion and consistence. The textural classes used for LRI were used to classify and a surface soil texture map was generated. The area extent and their geographical distribution in the subwatershed is shown in Fig. 9. An area of about 840 ha (18.13%) has a clay texture at the surface. The most productive lands concerning surface soil texture are clayey soils (93.19%) with an infiltration rate of 2mm/hr that have a high potential for retention and availability of

water and nutrients but, have more problems with drainage, infiltration, workability and other physical problems. The sandy loam (1182 ha), clay loam (307 ha), sandy clay (634 ha) and sandy clay loam (1570 ha) texture classes occupied the majority of the sub-watershed area. the soils of watershed viz., sandy clay loam (10 mm/hr), sandy clay (14 mm/hr) and sandy loam (28 mm/hr) resulted in more infiltration rate than clay (2 mm/hr) and clay loam soils (7 mm/hr). The surface soil textural class provides a guide understanding soil-water retention and to availability, nutrient holding capacity, infiltration, workability, drainage, physical and chemical behaviour, microbial activity and crop suitability [14-16].



Fig. 7. Seasonal Budyko curve of Artal sub-watershed of Karnataka (2016-2021)



Fig. 8. AET/P Ratio of Artal sub-watershed of Karnataka



Fig. 9. Surface soil texture and infiltration rate map of Artal sub-watershed of Karnataka



Fig. 10. Pre- monsoon Ground water depth of Artal sub-watershed of Karnataka



Fig. 11. Post-monsoon Groundwater depth of Artal sub-watershed of Karnataka



Fig. 12. Soil and Water Conservation Plan map of Artal sub-watershed of Karnataka

3.7 Ground Water Depth Measurement and Kriging

Groundwater levels fluctuate naturally in response to a sequence of climatic events and to constraints imposed by hydrogeologic and topographic characteristics. The groundwater level is influenced by borewell recharge, discharge, topography of land, soil texture etc. Trend analysis of water table depths indicates marked spatial variations of groundwater levels in the Artal sub-watershed of the study area. The mean depth of groundwater observed from ground level during pre and post-monsoon. During pre and monsoon the average ground water depth highest of 46-50 mbgl and lowest of 5-10 mbgl. These data indicate marked spatial variability in the distribution of wells with distinct rates of change across the different geomorphic units visible [17]. The groundwater resource of a region is one of the building blocks for the balanced economic development of the area. The water table represents the groundwater reservoir, and changes in its level represent the changes in groundwater storage [18].

The map of groundwater elevation determined by this method is presented in Figs. 10 and 11. It

shows that the highest groundwater elevation occurred in the west to the western part of the study area and the lowest groundwater elevation was obtained in the northeastern part of the study area. The groundwater elevation gradients are higher in the northern part and gradually decrease towards the southern parts and the general flow occurs from north to south [19]. "The groundwater table is deep on the upstream side and shallow on the mid and valley sides. This is possibly due to the flux that the water drains downslope to bring the soil moisture to the field capacity" [20]. "In addition, the soil depth on the upslope is shallow, which means it dries out faster than the deep soils due to evaporation; therefore, the quantity of water flowing toward the well is declining faster as compared with the well in the deep soil" [21].

3.8 Soil and Water Conservation Treatment Plan

For preparing the soil and water conservation treatment plan for the Artal sub-watershed, the land resource inventory database was generated and transformed as information through a series of interpretative (Thematic) maps using a soil phase map as a base. A map showing soil and water conservation plan with different kinds of structures recommended has been generated which shows the spatial distribution and extent of area. A maximum area of about 875 ha (18.9%) requires graded bunding and a 3657 ha (78.96%) area requires contour bunding (Fig. 12).

4. CONCLUSION

The total geographical area of the Artal subwatershed is 4632.04 ha and the sub-watershed is primarily composed of loamy sand followed by sandy clay loam, sandy loam and clay. The infiltration rates within the sub-watershed varied based on the type of soil and ranged from 3 to 28 mm/hr. In sub-watershed the mean annual ET is lower than the mean annual rainfall. This means that in the annual time scale of the water budget demand side is lower than the supply side. The focus in this sub-watershed should be to develop an improved water use efficiency (WUE) in agriculture and hence efforts should focus on the demand side of water use. The groundwater map shows the natural topography and prevailing conditions in the watershed are favorable for declining water table. The point recharge and farm ponds may be constructed in the lowermost corner of the agricultural fields to increase the natural recharge of rainwater during the monsoon period.

DISCLAIMER (ARTIFICIAL INTELLIGENCE)

Author(s) hereby declares that NO generative AI technologies such as Large Language Models (ChatGPT, COPILOT, etc) and text-to-image generators have been used during writing or editing of this manuscript.

COMPETING INTERESTS

Authors have declared that no competing interests exist.

REFERENCES

- Sadeghi SHR. Watershed Management in the 21st Century. In: Yousuf, A. and Singh, MJ. (eds.) Watershed Hydrology, Management and Modelling. CRC Press, Taylor and Francis, Boca Raton. 2020: 152-162.
- 2. Anonymous. A source book for soil and water conservation measures. Foundation for Ecological Security, Anand, Gujarat. 2018:1-239.

- Wurbs R. Dissemination of generalized water resources models in the United States. Water International.1998;23: 190-198.
- 4. Mustafaa JS, Mawlood DK. Mapping groundwater levels in Erbil Basin, American Academic Scientific Research Journal for Engineering, Technology and Sciences. 2023;93(1):21-38.
- 5. Varouchakis EA, Hristopulos DT, Karatzas GP. Improving kriging of groundwater level data using nonlinear normalizing transformations - a field application, Hydrological Sciences Journal. 2012;57(7): 1404-1419.
- 6. Alexander Samuel, Atkilt Girma, Amanuel Zenebe, Tesfaalem G. Spatio-temporal variability of evapotranspiration and crop water requirement from space, Journal of Hydrology. 2018; 567:732-742.
- Melloulia HJ, Van Wesemaelb B, Poesenc J, Hartmann R. Evaporation losses from bare soils as influenced by cultivation techniques in semi-arid regions, Agricultural Water Management, 2000;42:355-369.
- Raghavendra Prasad K, Kantha Rao B, Patra GK, Himesh S, Sheshakumar G. Annual and seasonal trends in actual evapotranspiration over different meteorological sub-divisions in India using satellite-based data, Theoretical and Applied Climatology. 2023; 52:999–1017.
- 9. Dey P, Mishra A. Separating the impacts of climate change and human activities on streamflow: A review of methodologies and critical assumptions, Journal of Hydrology. 2017; 548:278–290.
- 10. Yang D, Sun F, Liu Z, Cong Z, Lei Z. Interpreting the complementary relationship in non -humid environments based on the Budyko and Penman hypotheses, Geophysical Research Letters. 2006; 33:1-5.
- 11. Kyeung Kim, Hakkwan Kim, Hyunji Lee, Sang-Min Jun, Soonho Hwang, Jung-Hun Song and Moon-Seong Kang, Development and Assessment of Watershed Management Indicators Using the Budyko Frame work Parameter, Sustainability. 2021; 13:1-15.
- Gerrits AMJ, Savenije HHG, Veling EJM, Pfister L. Analytical derivation of the Budyko curve based on rainfall characteristics and a simple evaporation model. Water Resources Research. 2009; 45:1-15.

- Zhao J, Wang D, Yang H, Sivapalan M. Unifying catchment water balance models for different time scales through the maximum entropy production principle. Water Resources Research. 2016; 52 :7503–7512.
- 14. Christelle Basset, Majdi Abou Najm, Teamrat G, Xiaoxiao Hao, Andr'e D. How does soil structure affect water infiltration? A meta-data systematic review, Soil and Tillage Research. 2023; 226:1-15.
- 15. Nawaz MF, Bourri'e G, Trolard F. Soil compaction impact and modelling. a review. Agron. Sustain. Dev. 2013; 33 (2):291–309.
- 16. Huang J, Hartemink AE. Soil and environmental issues in sandy soils. Earth Sci. Rev. 2020; 208;103295.
- 17. Suneel Kumar Joshi, Sanjeev Gupta, Rajiv Sinha, Alexander LD, Shive Prakash R, Shashank S, Philippa JM, Van Dijk WM. Strongly heterogeneous patterns of

groundwater depletion in Northwestern India, Journal of Hydrology.2021; 598:126-492.

- Raghavendra G. Groundwater fluctuation and the flow pattern in the Kamarwadi subbasin of Bhima River. Res Rev J Eng Technol. 2013;73:101-108.
- Nikroo L, Zare MK, Sepaskhah AR, Shamsi SRF. Groundwater depth and elevation interpolation by kriging methods in Mohr Basin of Fars province in Iran. Environmental Monitoring and Assessment. 2010; z166:387-407.
- 20. Addisie MB. Groundwater recharge estimation using water table fluctuation and empirical methods. H₂Open Journal. 2022;5(3):457-468.
- 21. Cholo BE and Tolossa JG. Identification of groundwater recharge and fow processes inferred from stable water isotopes and hydraulic data in Bilate river watershed, Ethiopia, Hydrogeology Journal. 2023; 31: 2307-2321.

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